

Recent fog forecasting research at the Met Office

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Outline

- Why fog?
- Fog related research
 - LANFEX
 - Higher resolution fog modelling
- Future plans
- Particular acknowledgements to Ian Boutle



- Significant high impact weather, particularly for transport
- Every year fog causes serious road traffic accidents
- Fog at airports leads to reduced landing and take off rates \rightarrow flight delays
- At Heathrow, fog \rightarrow flight cancellations and diversions
 - costs airlines and BAA £millions
 - the Met Office supplies weather services to Heathrow



What makes fog hard to forecast?

Met Office

- Several "NWP Problem Group" tickets relate to poor fog forecasts
- Physically complicated
 - Result of feedbacks and imbalances between many small scale processes
- Spatial distribution subject to both large scale and local influences
 - Both scales introduce uncertainty







 Initial formation and evolution strongly dependent on fine imbalances between surface, turbulent, radiative and microphysical processes

- Implies spatially highly variable
- Variability further enhanced by a positive feedback via longwave radiative cooling
 - Potential to transition to "mature", well-mixed fog layer





Met Office forecast models

- Global: deterministic N768 (~17km) L70
 - MOGREPS-G= 33km global ensemble, 12 members
- UK: deterministic ("UKV") 1.5km, L70
 - MOGREPS-UK = 2.2km UK ensemble , 12 members
- London Model, 333m, L70
 - "Downscales" UKV



Quantifying uncertainty

• MOGREPS-UK samples a range of plausible solutions

- 12 members at 2.2km resolution over the UK
- Currently differ only in their initial conditions and boundary forcing (both direct from global ensemble)
 - So only samples uncertainty in the large scale state of the atmosphere
- Additional uncertainty should come from variability of small scale processes
 - random parameters to be introduced this winter
- Useful forecast tool for quantifying probability of an event
- Also useful for identifying systematic errors in model physics



Met Office MOGREPS-UK visibility postage stamps, T+9

06Z 15th Nov 2012

/is<50 m 50 m< vis< 100 m 100 m< vis< 200 m 1 km< vis< 5 km vis> 10 km No visibility data



AVABZ Atmos vis at 1.5m (incl precip) m at -1.000 metres At 062 an 15/11/2012, from 21Z on 14/11/2012









AAAB2 Atmos vie at 1.5m (Incl precip) m at -1.000 metree At 062 on 15/11/2012, from 212 on 14/11/2012











AAABZ Atmos vis at 1.5m (incl. precip) m at -1.000 metres At 08Z on 15/11/2012, from 21Z on 14/11/2012



1.000 metres



AAABZ Atmos vis at 1.5m (incl precip) m at -1.000 metres At 062 nn 15/11/2012, fram 212 nn 14/11/2012

AAABZ Atmas vis at 1.5m (incl precip) m at -1.000 metres At O5Z on 15/11/2012, from 21Z on 14/11/2012

AvABZ Atmos vis at 1.5m (incl precip) m at -1.000 metres Av 05Z on 15/11/2012, from 21Z on 14/11/2012









Large-scale differences in low cloud advection dictates location of fog in SE England

AAABZ Atmos vis at 1.5m (incl precip) m at -1.000 metres At 08Z on 15/11/2012, from 21Z on 14/11/2012



AAABZ Atmos vis at 1.5m (incl precip) m at -1.000 metr At 06Z an 15/11/2012, from 21Z an 14/11/2012





AAAB7 Atmos ice cloud amount At 052 on 15/11/2012, from 212 on 14/11/2012





AAABZ Atmose low cloud amount At G6Z on 15/11/2012, from 21Z on 14/11/2012







Met Office

Convective-scale (<4km) Met UM physical parametrizations relevant to fog

- Scale-aware blending of boundary-layer and Smagorinsky turbulence schemes
 - Gives a scale-dependent blend as the flow transitions from unresolved to resolved turbulence
 - Self-adapting for all high resolution configurations
- Scale-aware warm rain microphysics
 - Directly represents an appropriate subgrid variability in the warm rain microphysical conversion rates
- "murk" = single variable for total aerosol mass = air mass characteristics
 - Used to diagnose visibility and thence assimilated
 - Has surface sources, is advected by dynamics, mixed by turbulence, rained out,...
 - Sets cloud droplet number for microphysics
 - ...and for radiation, starting this winter (but murk has no direct radiative effect)
- Diagnostic **cloud** scheme with specified distribution width (RH_{crit})
- JULES 9-tile **surface** scheme
- Radiation: 2-stream (Edwards & Slingo), accounting for terrain slope
- No convection scheme

Fog research



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Recent fog studies at the Met Office

Met Office

- Ongoing projects since 2006 (e.g. Price 2011, Porson et al. 2011, Price et al 2015)
- Observations in radiation fogs, adiabatic fogs, stratus fogs at Met Office Cardington site
- Combined with modelling studies LES, SCM and NWP (both deterministic and ensemble)
- Some key observations:-
 - Fog may not form as expected under sustained periods of very high RH (>98%)
 - Suspected importance of (unknown) local heterogeneity
 - Systematic bias in MetUM ensemble to lift fog into low stratus



Science questions

that a focussed field campaign might address

- Heterogeneity: why does fog form in one place and not another?
 - Dynamics (sheltering, cold pools, local turbulence, etc)
 - Thermodynamics (RH of air vs dew competition)
 - Aerosol (high RH but weak activation)
- Understanding fog evolution/propagation
 - Advection or propagation? Role of drainage flows
 - What are the controls on optical thickness and the transition to mature fog?
- Predictability
 - How robust is the heterogeneity?
 - What impact does large scale uncertainty play?
 - Can we quantify local uncertainty?
- Basic MetUM validation and informing parametrization development

COLPEX→LANFEX (2009-2010) (2014-2016)



Met Office Local And Non-local Fog EXperiment

- 18 month campaign to examine development and evolution of (primarily) radiation fogs (Autumn 2014-Spring 2016).
 - Deploying longterm networks of instruments (flux towers, surface sites, dopler lidar, etc)
 - IOPs with sondes, tethered balloon
- High resolution modelling run in parallel
- Two sites: Met Office Cardington and Shropshire hills





100 200 300 400 500



Initial modelling studies

- Running high resolution (from 1.5km to 100m or finer) MetUM case studies
 - Identify systematic errors (use of ensembles)
 - Sensitivity tests and coarse-graining to inform parametrization developments
- Examples of recent results, focussing on each parametrization in turn
 - Cloud scheme
 - Orography
 - Stable boundary layer and surface coupling
 - Aerosols and microphysics

The role of the cloud scheme



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The London Model (LM)

Met Office

333m grid-length MetUM nested inside the UKV (1.5km)

Better representation of terrain and land surface characteristics

300x200 grid-points (100x66km domain)

Running (at 6Z and 21Z) every day since 2014



London Model domain



- Parametrizations are the same as UKV except for the cloud scheme
 - diagnostic scheme with specified distribution width (RH_{crit})
 - increase RH_{crit} in LM, consistent with its smaller grid-boxes

Case Study 18/10/2013



- Patchy fog in vicinity of Heathrow (H), fog observed at Northolt (N)
- Fog erroneously absent in UKV, but present in LM
- LM had breaks in a thin low cloud layer earlier in the night where UKV didn't
 - Allowed surface temperature to cool, and fog to form



Extra LM moisture variability



- Calculate resolved variability in RH in a 5x5 grid-point (1.67km) region of the LM
- Equivalent to the additional variability the UKV should need to parametrize
- Additional resolved variability is significantly less than parametrized difference in RH_{crit}
- Suggests RH_{crit} should vary less with resolution
- Also shows one single number for RH_{crit} doesn't exist – need a more sophisticated cloud scheme



Forecast time

The role of small scale orography



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How does the orography affect fog?

Met Office

 Calculate effective stability function (McCabe & Brown 2007)

$$f_{het}\left(\left\langle Ri\right\rangle\right) = \frac{wb_{het}}{wb_{1.67}} f\left(\left\langle Ri\right\rangle\right)$$

- wb=buoyancy flux, Ri=Richardson number
- Quantifies the additional mixing created at 1.67km (5x5 grid points) due to resolved variability in LM
- UKV uses the "sharp-tail" to parametrize stable BL mixing
- More turbulent mixing in LM, less fog





- Lots of additional variability in the near surface windspeed seen in LM
 - Closely tied to variability in the orography

The role of stable boundary layers and surface coupling





UK temperatures Diurnal cycle verification for June 2014

- Suppressed diurnal cycle (UKV and GM) in near-surface temperature
 - Systematic physics problem
- Biases amplified when sampled over clear sky cases
 - Not simply excessive cloud
- Warm bias at night
 - Excessive SBL mixing?
- Cold bias by day
 - BL depth/entrainment wrong?
 - Surface fluxes/Bowen ratio wrong?
 - Cold bias above SBL?
- Excessive coupling between atmosphere and soil could explain both





Comparisons at Cardington

- Off-line JULES (driven by observed T, q, U and down-welling radiation)
 - Reproduces suppressed diurnal range
 - Shows strong sensitivity to land surface properties
- Top soil level (10cm thick) in UKV has greater diurnal range than observations at 1cm
 - · Alleviated by reducing the fraction of bare soil





Clear sky sensitivity to SBL mixing Cardington clear-sky case study, 16th April 2014

- Reduced mixing in stable boundary layers
 - Minimum mixing length (5m cf 40m)
 - Sharper than "Sharpest" stability function
 - Use surface T in lowest-level Ri
- Improves near-surface T profile
- Enhanced vertical resolution shows little impact





LANFEX case study

visi screen

14:00 16:00 18:00 20:00 22:00 00:00 02:00 04:00 06:00 08:00 10:00

Time

Visibility

Met Office

- More realistic grass surface and reduced SBL mixing, "Real Grass", improves surface temperature and fog formation
- But soon after fog forms the model's becomes optically thick and spuriously warms the surface

10⁵

10⁴

10³

10²

10

Visibility in air / m



20

14:00 16:00 18:00 20:00 22:00

00:00 02:00 04:00 06:00 08:00 10:00

Time

The role of aerosol and cloud microphysics



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Slowing the transition to optically thick fog

- Weak cooling rates in fog imply limited droplet activation
 - Hence already "taper" to small N at surface
- Impose N=50cm⁻³ below 50m
 - Reduces fog optical thickness giving cooler surface
 - Surface fog deposition rate still realistic
- Still lacking dew deposition in the early evening despite reasonable surface T



14:00 16:00 18:00 20:00 22:00 00:00 02:00 04:00 06:00 08:00 10:00

Time





Slowing the transition to optically thick fog

 Optically thinner fog also improves temperature profile and heat fluxes

120

100

PS36

Real Grass New Tape



Sens heat flux / W m-2

Office

03:30

Temperature

Potential

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- Fog is hard!
- Large-scale meteorology is a key uncertainty
 - Use of ensembles
- Highlights weaknesses in parametrizations
 - Result of feedbacks and imbalances between many small scale processes



- Met Office fog research is focussed around LANFEX
- On-going work is investigating:
 - Cloud scheme: needs to be more adaptive
 - PC2, parametrize RH_{crit}?
 - Turbulent mixing in stable BL: still too strong?
 - May also need to represent effect of small hills properly (how?)
 - Surface characteristics: critical for accurate diurnal T
 - Need much better ancillary information (vegetation coverage, canopy height, LAI, etc)
 - Aerosol ↔ droplets: important for accurate microphysical properties (for radiation and settling) and thence fog evolution
 - How complex?
 - Tweak taper profile; prognostic cloud drop number; full chemistry?

Thank You

Questions?



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5% error at 1000 km = 100% error at 50 km



Stratus lowering to fog

Met Office

- Simpler system, well-mixed boundary layer, thicker cloud (better resolved)
- But cloud base height very sensitive to the boundary layer moisture and temperature budgets
 - Lifting base: warming from surface, entrainment drying, precip to surface
 - Lowering base: LW cooling, cooling/moistening from surface, evaporation of precip





LANFEX modelling domains

Met Office • Com

Comparing grids from 1.5km down to 100m (or less)

Shropshire

Surface altitude





Cardington

Surface altitude





Excessive cloud in UKV



Forecast time



What works for 1 case... 23rd Sept 2013, more low cloud, less fog



- Cloud scheme dominates impact
- Here, the UKV RH_{crit} value works better – LM was too foggy due to lack of cloud
- Hence can't simply tune RH_{crit}, need adaptivity
 - PC2? Parametrize RH_{crit}?





Dew deposition

Met Office • There is enormous variability in dew deposition at Cardington (Met Office Research Unit)

- By location around the site (see below)
- By surface characteristics (eg. long/short/dead grass)





Sussacta cae texe texe regainse i tyla setimete in (\$7(20)30)

 Trees/hedges 2-3K warmer than fields so gridbox mean T will be biased warm compared to grass obs





Verification of RH > 96%

Met Office

PS31: revised stable turbulence

- Getting near-surface air close to saturation has got to be a prerequisite for good fog forecasts
- UKV got a lot better at predicting RH close to saturation (at least in winter) from 2013 (PS31)
 - deliberate anticipated impact of changes to stable boundary layer mixing
- But what about actual fog forecasts?





Verification of fog (vis<1km)

UK-UKV - T+27

- UKV fog frequency bias is almost always very low
 - Even higher frequency biases are in winter • and when there is relatively little fog
 - So statistical significance isn't high
 - But also partly due to missing low cloud
- Why is fog frequency bias much worse (lower) than that for high RH?
 - Often get high RH without fog
 - E.g., competition between dew deposition and fog



UK-GM - T+24

→ UK-E4 - T+24



Fog is also patchy

- ETS for fog (vis<1km) is poor (UKV~0.1)
- But getting a hit at a site for an infrequent and patchy quantity like fog is hard
- Starting to use neighbourhoods to give a local fog probability
 - UKV scores show some skill even at 4km scale (3x3 points)
 - MOGREPS-UK ensemble gives significantly better skill than UKV

